

# 桐柏高压变质地体：对桐柏-大别-苏鲁高压/超高压变质带构造框架和俯冲/折返机制的制约\*

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**Abstract** The Tongbai orogen is situated at the junction between the Qinling and Dabie-Sulu orogenes; hence it may hold a key to understanding the geological relationships and tectonic evolution of different terranes in the Qinling-Tongbai-Dabie-Sulu orogenic belt. The Tongbai high-pressure (HP) metamorphic terrane comprises two HP slices (I and II) and a tectonic mélange zone in the north and a blueschist-greenschist zone in the south. The HP slice I is represented by the northern and southern eclogite zones occurring on the two sides of the Tongbaishan antiform. The peak metamorphic conditions are estimated to be 530~610°C and 1.7~2.0GPa for the northern zone and 460~560°C and 1.3~1.9GPa for the southern zone. The HP slice II is represented by the metamorphic enclaves in gneissic granites in the Tongbai complex. Their peak metamorphic conditions are inferred to be within the eclogites facies field of <700°C and >1.2GPa, and retrograde P-T conditions at 660~700°C and 0.80~1.03GPa. U-Pb, Lu-Hf, Rb-Sr and Ar-Ar multichronometric data indicate the peak metamorphism of the HP slice I taking place at 255Ma, whereas cooling to the closure temperature of muscovite at 238Ma. By contrast, the metamorphic ages of the HP slice II are as young as 232~220Ma, whereas gneissic granites dominating the Tongbai complex were emplaced at 140Ma. This suggests that, when the HP slice II was subducted to deep levels and underwent HP metamorphism, the HP slice II overlain it might have been exhumed to mid- to upper-crustal levels. This testifies a diachronous (differential) subduction/exhumation model of different HP and ultrahigh-pressure (UHP) slices proposed for the western Dabie, eastern Dabie and Sulu regions. Furthermore, the above data demonstrate the commencement of subduction of the South China continental crust at the Late Permian, which also represents the transitional time of oceanic to continental subduction between the North and South China blocks. Based on the similarity between the Tongbai and North Dabie complexes, we consider that a lower P-T environment and the absence of UHP rocks in the Tongbai HP metamorphic terrane might be resulted from the shallower subduction of the South China block towards the west. The final west-narrowing architecture of the Tongbai-Dabie HP/UHP metamorphic belt is related to the Early Cretaceous extrusion towards the east.

**Key words** HP/UHP; Diachronous subduction/exhumation; Tectonic architecture; Tongbai-Dabie-Sulu orogen

**摘要** 桐柏造山带位于秦岭造山带和大别-苏鲁造山带之间,是揭示秦岭-桐柏-大别-苏鲁巨型造山带中各地质体之间构造关系及地质演化差异的关键地区。桐柏高压变质地体主要由两个高压岩片(I和II)及其北侧的构造混杂岩带和南侧的蓝

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片岩-绿片岩带构成。高压岩片 I 以北、南两条榴辉岩带为代表, 构成桐柏山背形构造的两翼, 其峰期变质条件分别为 530~610°C、1.7~2.0GPa 和 460~560°C、1.3~1.9GPa。高压岩片 II 以桐柏杂岩中的变质岩包体为代表, 其峰期变质条件推测在 <700°C、>1.2GPa 的榴辉岩相范围内, 而退变质条件为 660~700°C、0.80~1.03GPa。U-Pb、Lu-Hf、Rb-Sr 和 Ar-Ar 同位素年代学研究表明, 高压岩片 I 的峰期变质时代为 255Ma, 冷却至白云母封闭温度的时代为 238Ma; 而高压岩片 II 的主期变质作用发生在 232~220Ma, 作为桐柏杂岩主体的片麻状花岗岩则侵位于 140Ma。这说明, 高压岩片 I 和 II 分属于两个时代不同的俯冲/折返岩片, 当高压岩片 II 被俯冲到地壳深处并经受高压变质时, 其上覆的高压岩片 I 已经折返到中上地壳的水平。这一结果验证了在西大别、东大别和苏鲁地区提出的高压/超高压岩石的穿时(或差异)俯冲/折返模型, 同时说明华南大陆地壳最早的大别造山带发生在晚二叠世, 这也代表华北与华南陆块之间从洋壳俯冲转化为陆壳俯冲的时间。基于桐柏杂岩与北大别杂岩的可比性, 认为桐柏高压变质地体相对低温低压的变质环境以及超高压岩石的缺乏缘于华南陆块的俯冲深度向西逐渐变浅, 而早白垩世的构造挤压造成了桐柏-大别高压/超高压变质带东宽西窄的构造格局。

**关键词** 高压/超高压; 穿时俯冲/折返; 构造框架; 桐柏-大别-苏鲁造山带

**中图法分类号** P545

## 1 引言

秦岭-桐柏-大别-苏鲁造山带代表华北与华南陆块之间的碰撞带, 其内产有世界上出露规模最大的高压/超高压变质带。该造山带被一系列北北东向断裂或盆地切割为五个主要构造块体, 从西向东分别为秦岭、桐柏、西大别、东大别和苏鲁地体(图 1)。从区域上看, 以含有柯石英为特征的超高压变质岩主要分布于东大别和苏鲁地区, 西大别地区则以高压变质杂岩为主, 柯石英仅发现于新县穹窿核部的榴辉岩及其伴生岩石(李学燮和徐培仓, 1993; Zhang and Liou, 1994; Cui and Wang, 1995; Liu JB et al., 1998; Liu XC et al., 2004a)。进一步向西, 在桐柏地区仍有高压榴辉岩出露

(魏春景等, 1999; 索书田等, 2001a; 刘晓春等, 2005), 但秦岭地区(不包括古生代高压/超高压变质岩)则只有低温高压蓝片岩的报道(周高志等, 1991)。高压/超高压变质岩沿造山带的规律性分布似将有助于阐明华北与华南陆块之间的构造演化过程和大规模超高压变质作用的发生机理。

桐柏地区位于秦岭-桐柏-大别-苏鲁造山带的中部, 向西越过南阳盆地与秦岭造山带相接, 向东以大悟断裂为界与大别-苏鲁造山带相连, 因而是揭示秦岭造山带与大别-苏鲁造山带的构造关系及地质演化差异的关键地区。然而, 与研究程度较高的秦岭造山带和大别-苏鲁造山带相比, 对桐柏造山带的研究还相当薄弱, 特别是桐柏杂岩的构造性质及其与两侧高压榴辉岩带的关系尚不十分清楚, 该区是否有超高压变质作用发生也缺乏可靠的证据和明确的结论。近年来, 我们结合国家和部门的科研项目, 对桐柏造山带进行了详细的

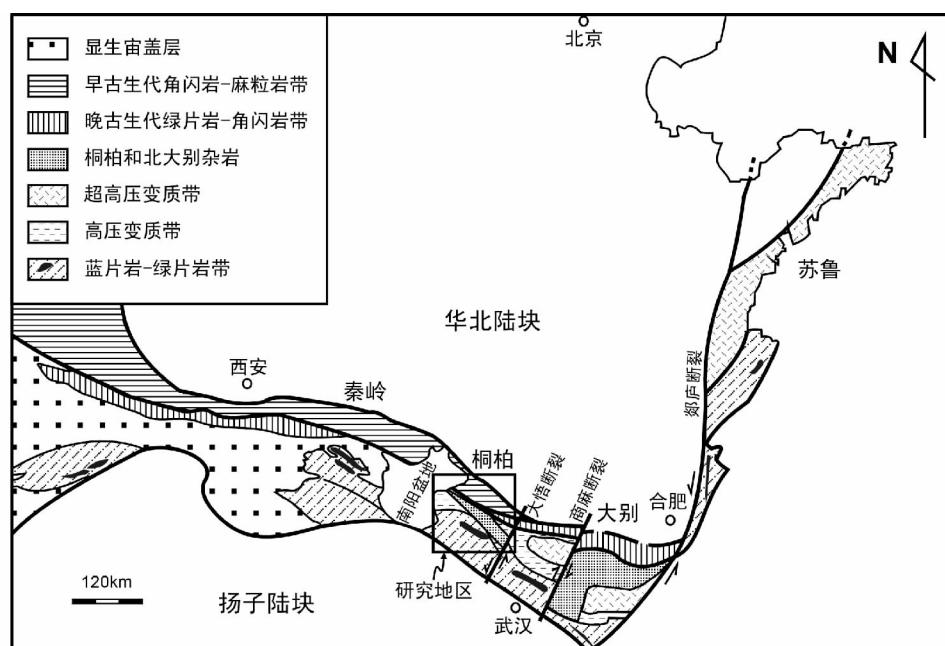


图 1 秦岭-桐柏-大别-苏鲁造山带区域构造图(据 Liou et al., 1996 修改)

Fig. 1 Regional tectonic map of the Qinling-Tongbai-Dabie-Sulu orogenic belt (modified after Liou et al., 1996)

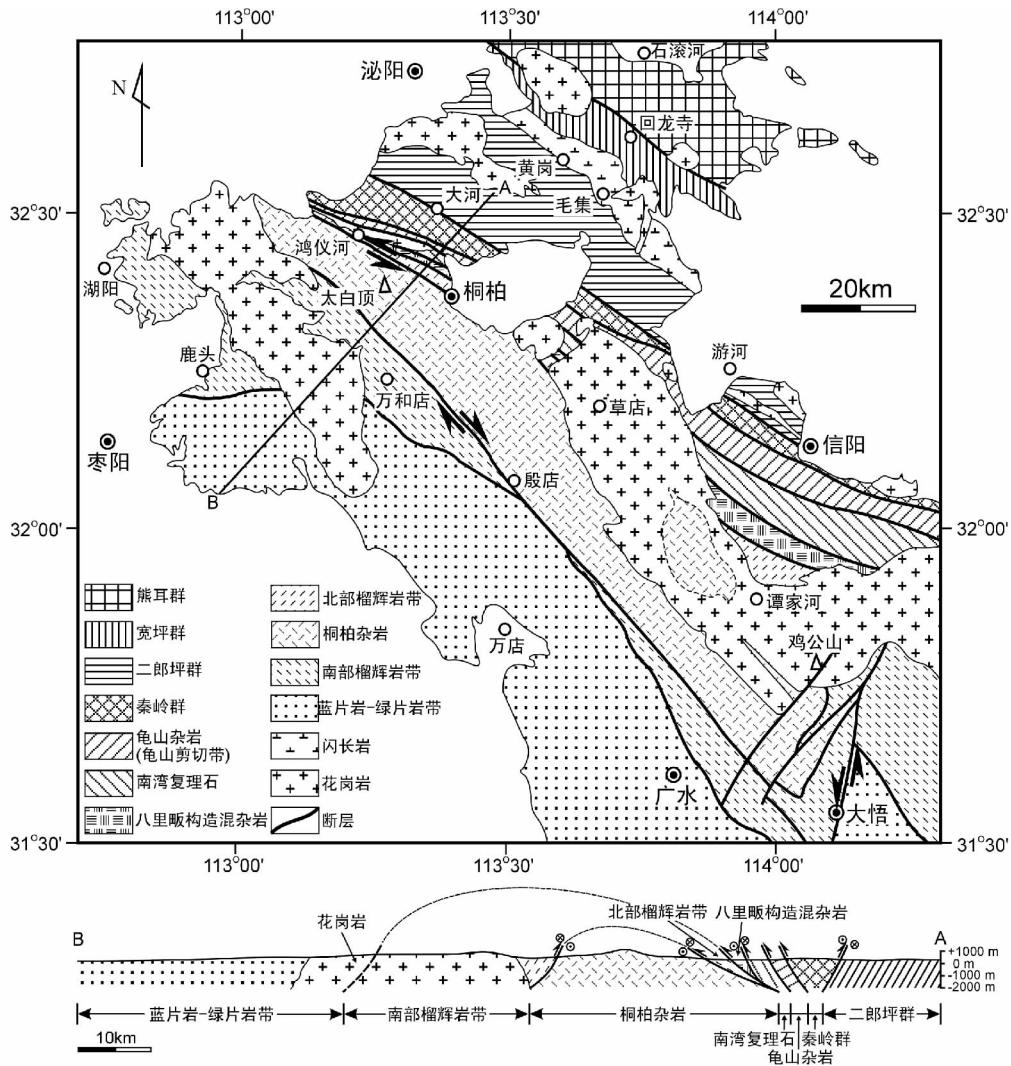


图2 桐柏地区地质简图及剖面图

Fig. 2 Simplified geological map and a generalized profile in the Tongbai area

野外地质调查和构造地质学、岩石学、地球化学及同位素年代学综合研究,基本上厘定了桐柏造山带的构造框架和演化过程。本文系统总结和评述了对桐柏高压变质地体最新的岩石学和年代学研究成果,并试图通过区域对比来探讨桐柏-大别-苏鲁高压/超高压变质带的结构以及大陆深俯冲和折返的过程。

## 2 区域地质背景

桐柏地区主要由两套造山系统构成,北侧为古生代增生造山系统,南侧为中生代碰撞造山系统(图2)。古生代增生造山系统包括宽坪群、二郎坪群、秦岭群、龟山杂岩(龟山剪切带)和南湾复理石。除龟山杂岩外,这些岩石构造单元与秦岭造山带中的相似岩群可一一对应(张国伟等,1988; Xue et al., 1996; Meng and Zhang, 2000; Ratschbacher et al., 2003)。中生代碰撞造山系统则包括八里畈构造混杂岩带

(定远组)、北部榴辉岩带(浒湾组)、桐柏杂岩、南部榴辉岩带和蓝片岩-绿片岩带(刘晓春等,2005; Liu XC et al., 2008, 2010)。这些岩石构造单元构成了桐柏山背形的核部和两翼,基本上可以与西大别高压/超高压变质地体相对比(Liu XC et al., 2004a, b),其最重要的差别在于核部桐柏杂岩的产出以及超高压变质单元的缺乏。各岩石构造单元主要表现为一系列经历了复杂构造变形的大型透镜状构造岩片,边界多为早期逆冲、后期伸展所形成的韧性剪切带,少数叠加了脆性断裂(Liu X et al., 2011),两套造山系统的边界即为华北-华南陆块最终碰撞的缝合线。早古生代的闪长质-花岗闪长质-花岗质深成岩侵入于北部二郎坪群中,早白垩世的花岗岩则广布于各岩石构造单元。

中生代碰撞造山系统中的八里畈构造混杂岩带位于北部榴辉岩带北侧,与南北两侧构造单元均为构造接触。主要由白(绢)云钠长片麻岩、绿片岩夹白云石英片岩组成,含有变质辉长岩和变质花岗闪长岩团块。变质作用仅为绿片岩

相,但变形强烈,除较大的变质辉长岩和花岗闪长岩团块外,其它岩石普遍经历了韧性剪切变形作用的改造,从而形成各种糜棱岩或糜棱岩化岩石。由于在西大别北部相同构造单元中获得的同位素年代学数据主要集中在新元古代,所以该带被认为是未被深俯冲的华南陆块残片(Hacker *et al.*, 2000; Liu XC *et al.*, 2004b; 刘贻灿等, 2006)。蓝片岩-绿片岩带位于南部榴辉岩带南侧,向东与西大别地区木兰山蓝片岩带相连。该带的层位性较好,下部主要由新元古代双峰式变质火山沉积岩系组成,上部则为白(绢)云母片岩、白云石英片岩与绿片岩互层,并构成一个宽缓的向斜。蓝片岩相变质作用的P-T条件为350°C、0.7GPa,并普遍存在由蓝片岩相向相对高温低压绿片岩相的转变(刘晓春等, 1989)。有关北、南榴辉岩带和桐柏杂岩的地质特征将在下文详细讨论,这里不再赘述。

### 3 北、南榴辉岩带(高压岩片 I)

在桐柏杂岩两侧出露北、南两条榴辉岩带,我们将其划归于高压岩片 I。北部榴辉岩带主要出露于鸿仪河-固庙-桐柏一带,向东延伸到谭家河北部,在大地构造部位上相当于西大别地区的浒湾组(浒湾剪切带),但缺少石炭纪榴辉岩。主要组成岩石为花岗质糜棱岩、白云钠长片麻岩、白云石英片岩和大理岩。榴辉岩及其退变产物石榴角闪岩以团块、透镜体或不规则条带产于白云钠长片麻岩和大理岩中,其中新鲜露头包括鸿仪河榴辉岩、桐柏榴辉岩和罗庄退变榴辉岩。南部榴辉岩带主要出露于七尖峰花岗岩体以西的黑龙镇-双河镇-大阜山一带,向东延伸到蔡家河,相当于西大别地区的红安群,主要由一套变质火山-沉积岩系组成,包括白云钠长片麻岩、白云石英片岩和大理岩。榴辉岩也以团块、透镜体或不规则条带产于白云钠长片麻岩和大理岩中,其中代表性的露头包括凤凰咀榴辉岩、大河口榴辉岩、上福田榴辉岩、程窑榴辉岩和大阜山榴辉岩等。值得一提的是,在南部榴辉岩带中还产出有大量的变质辉长岩和石榴角闪岩块体,以大阜山镁铁质-超镁铁质岩体为代表。岩体中的超镁铁质岩石保存较好,但镁铁质岩石只保留辉长岩结构,斜长石已全部转变为细粒黝帘石+钠云母+钠长石集合体,单斜辉石为角闪石+金红石所取代,而石榴石均呈冠状生长在角闪石的边缘(Liu XC *et al.*, 2008),其与榴辉岩的成因关系尚不清楚。

榴辉岩的典型矿物组合为石榴石+绿辉石+角闪石+多硅白云母+石英+金红石±绿帘石,部分样品遭受到绿帘角闪岩相退变质作用的改造。电子探针分析表明,所有榴辉岩中的石榴石都具有化学成分环带,从核部到边部主要表现为镁铝榴石分子(Prp)的升高和锰铝榴石分子(Sps)的降低,铁铝榴石(Alm)和钙铝榴石分子(Grs)则显示不规律的变化(刘晓春等, 2005),指示了石榴石的进变质生长。常规温压计算揭示,北带榴辉岩形成的P-T条件较高,为530~610°C、1.7~2.0GPa,南带榴辉岩较低,为460~560°C、1.3~1.9GPa

(刘晓春等, 2005; Liu XC *et al.*, 2008),这与使用THERMOCALC平均温压计算获得的结果(490~540°C、1.8~2.1GPa)基本一致(Cheng *et al.*, 2011)。考虑到两条榴辉岩带物质组成的相似性,推测二者可能代表同一个高压岩片的不同部位,分别对应于西大别地区的浒湾榴辉岩带(I带)和七角山榴辉岩带(V带)(Liu XC *et al.*, 2004a)。

我们对两条榴辉岩带中的典型榴辉岩和变质辉长岩体进行了SHRIMP锆石U-Pb定年(Liu XC *et al.*, 2008)(表1)。北部榴辉岩带中的罗庄退变榴辉岩含有变质成因的锆石,这些锆石具有典型的冷杉树状构造和极低的Th/U含量及Th/U比值,其内包裹石榴石、绿辉石、角闪石和绿帘石等榴辉岩相变质矿物,其中8个分析点给出加权平均年龄为255±6Ma,另外3个分析点给出的结果为216±14Ma。鸿仪河榴辉岩中的锆石均具有韵律环带,测试结果给出2组主要的年龄数据,其一是从768±16Ma到641±12Ma,其二为215±3Ma。南部榴辉岩带中大河口榴辉岩中的锆石主要有两种形态,其给出的年龄也不相同:细小的海绵状锆石颗粒年龄较老,且基本沿着不一致线分布,其上交点年龄为1645±15Ma,下交点年龄为257±16Ma;柱状具韵律环带锆石有2个颗粒,其中一颗给出的年龄为742±11Ma,另一颗为256~240Ma。程窑榴辉岩中的锆石较少,且给出的年龄结果比较分散,实际意义不大,但其全岩-石榴石-绿辉石-角闪石-多硅白云母Rb-Sr等时线年龄为253±11Ma,可以与变质锆石的年龄相对比。大阜山变质辉长岩中的锆石明显受到了晚期流体交代作用的影响,并形成两个年龄组,其加权平均年龄分别为662±10Ma和594±13Ma,应分别代表岩体侵位和流体改造的年龄。据此认为,桐柏地区北、南榴辉岩带的原岩形成于新元古代,而榴辉岩相变质作用发生在晚二叠世,约255Ma。

最近,Cheng *et al.* (2011)使用全岩-石榴石Lu-Hf定年方法对上述年代学结果进行了检验,获得北部榴辉岩带中鸿仪河附近一个样品的等时线年龄为246.9±3.2Ma,而南部榴辉岩带中双河镇北部两个样品的等时线年龄分别为252.3±3.4Ma和256.4±2.6Ma。鉴于石榴石中含有绿辉石包裹体,这3组等时线年龄均被解释为榴辉岩相变质年龄,进一步说明桐柏地区高压岩片 I 的变质时代应为晚二叠世。

与此同时,我们对榴辉岩的围岩-正片麻岩、副片麻岩和石英岩也进行了SHRIMP锆石U-Pb和白云母<sup>40</sup>Ar/<sup>39</sup>Ar定年(Liu XC *et al.*, 2008)。在这些样品中,所有的锆石均未发现二叠-三叠纪的变质生长边或者生长边很窄,其中1个正片麻岩中的锆石给出735±9Ma的岩浆结晶年龄,而3个变质沉积岩中的碎屑锆石揭示出3期变质年龄峰值,分别为2.49Ga、1.93Ga和1.85~1.82Ga,说明扬子陆块北缘曾遭受到古元古代强烈构造热事件的影响。白云母的<sup>40</sup>Ar/<sup>39</sup>Ar定年对桐柏地区高压变质作用的时代做出了进一步的限定,5个样品的分析结果分别为:鸿仪河正片麻岩238±2Ma,罗庄正片麻岩238±2Ma,大阜山石英岩234±2Ma,徐家岭正片麻

岩  $217 \pm 1$  Ma, 姚园副片麻岩  $334 \pm 4$  Ma。姚园副片麻岩中的白云母与钠云母连生,且本身具有较高的  $\text{Na}/(\text{Na} + \text{K})$  比值,因而推测其较老的年龄结果可能是由钠云母的混入或白云母中 Ar 的过剩造成的,不具有地质意义。

## 4 桐柏杂岩(高压岩片 II)

桐柏杂岩被断裂所围限,位于两条榴辉岩带之间,我们称之为高压岩片 II。桐柏杂岩主要由两部分组成,其主体是已强烈糜棱岩化的粗粒片麻状花岗岩(或称变形花岗岩或花岗质片麻岩),约占变形岩石总量的 80%。这套岩石经历了较高温度下的塑性流变、韧性剪切变形和走滑伸展,但除某些变形带中有少量白云母生成外,基本上没有发生明显的变质重结晶。Kröner *et al.* (1993) 使用单颗粒 Pb-Pb 蒸发技术获得其原岩侵位年龄为  $776 \pm 8$  Ma, 所以桐柏杂岩一直被认为是一个前寒武纪地质体。另一部分是这些片麻状花岗岩侵位时裹挟的大量变质岩包体,以细粒青灰色-灰黑色闪长质-奥长花岗质片麻岩为主,含少量副片麻岩、斜长角闪岩、大理岩和钙硅酸盐岩。变质岩包体常成群出现,有时也单独产出,多数为似层状、透镜状、串珠状和不规则状。包体大小悬殊,最大者宽度超过 1 km,最小者只有几厘米。这些包体的变质作用达角闪岩相,其变形面理与周围的片麻状花岗岩协调一致。

在鄂豫交界平靖关-武胜关这一狭长的高山区域内,我们发现了 4 个退变榴辉岩的原生露头和多个河谷转石,同时在多处发现了石榴角闪岩露头(崔建军等, 2009a; Liu XC *et al.*, 2010)。退变榴辉岩均呈透镜状或似层状产于闪长质-奥长花岗质片麻岩中。所有样品均已强烈退变,并明显分为两个粒级,粗粒矿物主要是石榴石和石英,细粒矿物包括成堆聚集的普通辉石+斜长石、角闪石+斜长石和个别黑云母+斜长石,推测他们分别是绿辉石、石榴石和多硅白云母的假象(Liu XC *et al.*, 2010)。石榴石的边缘均被细粒角闪石+斜长石集合体所取代,其内可包裹绿帘石、角闪石、斜长石、石英、榍石和钛铁矿。石榴石均具有化学成分环带,一般向边部 Prp 和 Alm 升高,Sps 降低,Grs 降低或近于恒定,显示了进变质的特征,但最外侧往往具有一个反向的扩散边。根据石榴石边部及邻近普通辉石、角闪石和斜长石成分获得退变质条件为  $660 \sim 700^\circ\text{C}$ 、 $0.80 \sim 1.03$  GPa,而由相关关系推测榴辉岩相的  $P$ - $T$  范围应位于  $<700^\circ\text{C}$ 、 $>1.2$  GPa 区域内(Liu XC *et al.*, 2010)。退变榴辉岩中石榴石显著的 Sps 环带、绿辉石假象复原成分中的低  $\text{Na}_2\text{O}$  含量( $<3\%$ ,但其成分可能受到后期重新平衡或交代作用的影响)以及锆石包裹体中柯石英的缺乏证明其从未达到超高压变质条件(Liu XC *et al.*, 2010)。由于退变榴辉岩与围岩紧密伴生,可能暗示桐柏杂岩中的变质岩包体都经受到榴辉岩相变质作用的影响。

我们对桐柏杂岩中不同类型的岩石进行了 SIMS 锆石 U-Pb 定年(Liu XC *et al.*, 2010)(表 1),基本上限定了桐柏杂

岩的地质演化历史。桐柏杂岩中少量副片麻岩包体起源于太古宙华南陆块上  $2960 \pm 9$  Ma 发生的岩浆活动,这些早期的岩浆岩在古元古代  $1966 \pm 22$  Ma 发生过变质作用。其它多数变质岩包体的原岩是罗迪尼亞超大陆聚合、解体过程中生成的岩浆岩,聚合岩浆作用发生在  $933 \pm 22$  Ma,裂解岩浆作用主要年龄段为  $863 \pm 11$  Ma 至  $742 \pm 30$  Ma。大多数变质岩包体记录了从  $232 \pm 6$  Ma 到  $220 \pm 3$  Ma 的三叠纪变质年龄,其中一个退变榴辉岩样品还具有一组  $208 \pm 3$  Ma 的年龄数据。桐柏杂岩的主体片麻状花岗岩侵位于早白垩世,两个样品的年龄分别为  $141 \pm 1$  Ma 和  $139 \pm 1$  Ma,其内可含有  $780 \sim 750$  Ma 的继承锆石残核,与 Kröner *et al.* (1993) 获得的单颗粒 Pb-Pb 蒸发年龄一致。同时,对一个侵入于变质岩包体中的同形伟晶岩脉也进行了定年,其侵位时间为  $131 \pm 1$  Ma(崔建军等, 2009b)。这些年代学结果说明,桐柏杂岩实际上是一个早白垩世花岗岩杂岩体,其中的变质岩包体则形成于三叠纪,而后二者又经历了统一的构造变形过程。所以,所谓的前寒武纪“桐柏杂岩”实际上并不存在。

桐柏杂岩的北、南两侧存在两条大型韧性剪切带,其运动学指向恰好相反,分别指示左行剪切和右行剪切(Webb *et al.*, 1999, 2001; 崔建军等, 2009b; 刘鑫等, 2010; Liu X *et al.*, 2011)。这两条韧性剪切带中的剪切面理分别朝 SSW 和 NNE 向倾斜,韧性剪切带及杂岩体内部发育的拉伸线理均稳定地朝 SEE 方向缓倾伏。而在杂岩体的东端发育低角度近南北向韧性剪切带,其糜棱面理和矿物拉伸线理均朝 SEE 方向低角度缓倾斜,运动学标志指示东侧地质体朝  $295^\circ \sim 310^\circ$  的方向逆冲(崔建军等, 2009b)。由不同岩石中角闪石、白云母和黑云母的  $^{40}\text{Ar}/^{39}\text{Ar}$  定年获得两组冷却年龄数据,其一为  $135 \sim 119$  Ma,其二为  $99 \sim 92$  Ma(Webb *et al.*, 1999, 2001; Cui *et al.*, 未发表数据),可能暗示桐柏杂岩在晚期经历了两阶段的抬升过程。

## 5 桐柏杂岩与北大别杂岩的对比

桐柏杂岩的岩石组合、结构样式以及变质变形序列与北大别杂岩非常相似(表 2)。北大别杂岩构成罗田穹窿的核部,其主体岩石是英云闪长质-奥长花岗质-花岗闪长质(TTG)片麻岩,其次为混合岩、斜长角闪岩、麻粒岩、大理岩和钙硅酸盐岩,局部含有榴辉岩透镜体。这些变质岩石被大量的花岗质岩石以及少量镁铁质-超镁铁质岩石所侵入(Zhang RY *et al.*, 1996; Jahn *et al.*, 1999; Tsai and Liou, 2000; Bryant *et al.*, 2004)。变质岩石记录了多期构造热事件,包括原岩形成于约  $800 \sim 750$  Ma,榴辉岩相变质作用及其随后的麻粒岩相/角闪岩相变质叠加发生于约  $230 \sim 210$  Ma,而最晚期的角闪岩相变质作用和部分熔融发生在约  $140 \sim 125$  Ma(Xue *et al.*, 1997; Hacker *et al.*, 1998, 2000; Bryant *et al.*, 2004; Xie *et al.*, 2004; Liu YC *et al.*, 2005, 2007a, b; Wu *et al.*, 2007; Zhao *et al.*, 2008)。变形和未变形花岗

表 1 桐柏高压变质地体现有同位素年龄统计

Table 1 Geochronological data in the Tongbai HP metamorphic terrane

构造单元	采样点	样品号	岩石类型	测年方法	年龄值(Ma)	地质解释	数据来源
北部榴辉岩带	桐柏	TB02-1	云母片岩	锆石 U-Pb	1820 ± 6	碎屑锆石继承年龄	[1]
	鸿仪河	HYH01-11	正片麻岩	锆石 U-Pb	735 ± 9	原岩年龄	[1]
	罗庄	GM02-1	退变榴辉岩	锆石 U-Pb	255 ± 6 216 ± 14	高压变质年龄 重结晶年龄	[1]
	鸿仪河	C100	榴辉岩	全岩-石榴石 Lu-Hf	246.9 ± 3.2	高压变质年龄	[2]
	鸿仪河	HYH1-3	榴辉岩	锆石 U-Pb	215 ± 3	重结晶年龄	[1]
	鸿仪河	HYH01-11	正片麻岩	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	238 ± 2	冷却年龄	[1]
	罗庄	GM02-4	正片麻岩	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	238 ± 2	冷却年龄	[1]
	徐家岭	GJZ10-1	正片麻岩	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	217 ± 1	冷却年龄	[1]
南部榴辉岩带	大阜山	DFS06-1	石英岩	锆石 U-Pb	2490 ± 11	碎屑锆石继承年龄	[1]
	姚园	SHZ03-5	副片麻岩	锆石 U-Pb	1846 ± 5	碎屑锆石继承年龄	[1]
	大阜山	DFS03-1	变质辉长岩	锆石 U-Pb	662 ± 10 594 ± 13	原岩侵位年龄 流体改造年龄	[1]
	大河口	SHZ08-1	榴辉岩	锆石 U-Pb	1645 ± 15 257 ± 16	碎屑锆石继承年龄 高压变质年龄	[1]
	双河镇北	C101	榴辉岩	全岩-石榴石 Lu-Hf	252.3 ± 3.4	高压变质年龄	[2]
	双河镇北	C102	榴辉岩	全岩-石榴石 Lu-Hf	256.4 ± 2.6	高压变质年龄	[2]
	姚园	SHZ03-5	副片麻岩	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	337 ± 2	Ar 过剩	[1]
	大阜山	DFS06-1	石英岩	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	234 ± 2	冷却年龄	[1]
桐柏杂岩	朱家沟	WHD04-2	副片麻岩	锆石 U-Pb	2960 ± 9 1966 ± 22	碎屑锆石继承年龄 碎屑锆石继承年龄	[3]
	曾家老门	PJG05-14	退变榴辉岩	锆石 U-Pb	863 ± 11	原岩侵位年龄	[3]
	水帘洞	T7	片麻状花岗岩	锆石 Pb-Pb	776 ± 8	侵位年龄	[4]
	水帘洞	T5	正片麻岩	锆石 Pb-Pb	746 ± 10	原岩侵位年龄	[4]
	廖庄	JTD003-2	石榴角闪岩	锆石 U-Pb	933 ± 12 232 ± 6	原岩侵位年龄 变质年龄	[3]
	平靖关	PJG01-8	石榴角闪岩	锆石 U-Pb	792 ± 11 225 ± 3	原岩侵位年龄 变质年龄	[3]
	火龙店	HLD09-2	闪长质片麻岩	锆石 U-Pb	748 ± 8 220 ± 3	原岩侵位年龄 变质年龄	[3]
	火龙店	HLD09-5	奥长花岗质片麻岩	锆石 U-Pb	742 ± 30 229 ± 4	原岩侵位年龄 变质年龄	[3]
	曾家老门	PJG05-2	退变榴辉岩	锆石 U-Pb	223 ± 3 208 ± 3	变质年龄 重结晶年龄(?)	[3]
	火龙店	HLD09-9	片麻状花岗岩	锆石 U-Pb	141 ± 1	侵位年龄	[3]
	熊堂沟	CZ007-1	片麻状花岗岩	锆石 U-Pb	139 ± 1	侵位年龄	[3]
	平靖关	PJG01-14	同变形伟晶岩	锆石 U-Pb	131 ± 1	侵位年龄	[5]
	鸡公山		花岗岩	锆石 U-Pb	128 ± 5	侵位年龄	[6]
	岩子河	D260B	韧性剪切带	白云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	131 ± 1	变形时代	[7]
	武胜关	D345A	韧性剪切带	黑云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	124 ± 1	变形时代	[7]
	广水	D347B	韧性剪切带	黑云母 <sup>40</sup> Ar/ <sup>39</sup> Ar	119 ± 1	变形时代	[7]

数据来源:[1] Liu XC et al. (2008); [2] Cheng et al. (2011); [3] Liu XC et al. (2010); [4] Kröner et al. (1993); [5] 崔建军等(2009b); [6] 李石和王彤(1991); [7] Webb et al. (1999)

岩以及镁铁质-超镁铁质岩石的侵位时间主要介于 135 ~ 117 Ma 之间(Xue et al., 1997; Jahn et al., 1999; Bryant et al., 2004; Zhao et al., 2005; Xie et al., 2006; Xu et al., 2007),部分变形花岗岩的侵入则发生在 143 ± 3 Ma 至 142 ± 3 Ma (Wang Q et al., 2007),与变质岩石的部分熔融(混合岩化)时间基本一致(139 ± 1 Ma; Wu et al., 2007)。与桐柏杂岩相似,北大别杂岩也显示了 NWW-SEE 方向的近水平伸展

构造(Hacker et al., 1998, 2000; Faure et al., 1999, 2003; Ratschbacher et al., 2000; Lin et al., 2009)。所以,北大别杂岩所记录的构造热事件似乎可以与桐柏杂岩一一对应。

关于北大别杂岩的构造演化尚存争议,这套杂岩是否也经历了超高压变质作用,还是只遭受到高压变质作用的影响也未有定论(Zhang RY et al., 2009)。实际上,比对于桐柏杂岩,北大别杂岩也可被看作是一个早白垩世片麻状(变形)

表 2 桐柏杂岩与北大别杂岩的相似性与差别

Table 2 Similarity and difference between the Tongbai and North Dabie complexes

对比	特征	桐柏杂岩	北大别杂岩
相似性	构造样式	桐柏山背形(穹隆),NWW-SEE伸展构造	罗田穹隆,NWW-SEE伸展构造
	岩石组合	变形花岗岩+变质岩(闪长质-奥长花岗质片麻岩及少量斜长角闪岩、副片麻岩、大理岩、钙硅酸岩岩和退变榴辉岩)	变形-未变形花岗岩+变质岩(TTG片麻岩、混合岩及少量斜长角闪岩、麻粒岩、石英岩、大理岩、钙硅酸岩岩和榴辉岩)
	变质/岩浆事件序列	(1)榴辉岩/角闪岩相变质232~220Ma;(2)变形花岗岩侵入140Ma;(3)未变形-弱变形花岗岩侵入128Ma	(1)榴辉岩/麻粒岩相变质230~210Ma;(2)角闪岩相变质、部分熔融和变形花岗岩侵入140Ma;(3)变形-未变形花岗岩及镁铁质-超镁铁质岩石侵入135~117Ma
差别	构造层位	相当于中地壳	相当于下地壳
	变质环境	相对低温低压(榴辉岩相变质<700℃、>1.2GPa;角闪岩相变质660~700℃、0.80~1.03GPa)	相对高温高压(榴辉岩相变质800~880℃、2.5GPa;麻粒岩相变质800~860℃、1.0~1.4GPa)
	早白垩纪事件	花岗质岩浆侵入	角闪岩相变质(680~780℃、0.5~0.7GPa);镁铁质-花岗质岩浆侵入

花岗岩地体,其内包裹大量(数量远多于桐柏杂岩)且规模较大的三叠纪高压或超高压变质岩石,只是二者之间的接触关系因后期强烈的变形作用、变质作用和部分熔融的改造而难以识别。然而,北大别杂岩明显处于相对高温高压的变质环境,其榴辉岩相变质幕P-T条件为800~880℃、2.5GPa,高压麻粒岩相变质幕为800~860℃、1.0~1.4GPa,角闪岩相变质幕为680~780℃、0.5~0.7GPa(Zhang RY et al., 1996; Tsai and Liou, 2000; Liu YC et al., 2007b)。此外,北大别杂岩上覆超高压岩片,而桐柏杂岩上覆高压岩片。Pb同位素地球化学研究表明,桐柏杂岩中变质岩包体的放射性Pb含量高于北大别杂岩,但低于南大别超高压岩片(张宏飞等,2001; Zhang HF et al., 2002, 2003, 2004)。如果象某些学者所建议的那样,北大别杂岩和南大别超高压岩片分别代表华南俯冲板片的下地壳和上地壳岩石(如Zhang HF et al. 2002; Li SG et al., 2005; Chen et al., 2006; Liu YC et al., 2007a),那么桐柏杂岩中的变质岩包体则相当于俯冲的中地壳岩石。

## 6 桐柏-大别-苏鲁高压/超高压变质带的构造框架

岩石学特征和温压计算结果表明,西大别地体超高压变质作用的峰期P-T条件低于东大别和苏鲁地体(Zhang and Liou, 1994; Eide and Liou, 2000; Liu XC et al., 2004a),桐柏杂岩中超高压榴辉岩的缺乏证明变质条件向西进一步降低,再向西越过南阳盆地(秦岭)则只出现低温高压的蓝片岩。绝大多数研究者推测,这种变质条件的规律性变化缘于造山带向西抬升幅度逐渐变小,致使超高压岩石在秦岭-桐柏地区未能出露于地表(Zhang and Liou, 1994; Zhang HF et al., 2004; Ernst et al., 2007)。根据可利用的地质和地球物

理资料(Hacker et al., 1996; Liou et al., 1996; Yuan et al., 2003; Dong et al., 2004, 2008; Liu XC et al., 2004a, b, 2010; Li SZ et al., 2010a),桐柏、西大别和东大别高压/超高压变质地体的三维地壳结构模型如图3所示。如果象Pb同位素成分所揭示的那样,桐柏杂岩中的变质岩包体代表俯冲的中地壳岩石,那么代表上地壳岩石的超高压岩片似应出现在桐柏杂岩和上覆的高压岩片I(即北、南榴辉岩带)之间,但事实并非如此。还有一种可能,即超高压岩片位于桐柏杂岩之下,但这样的地壳结构与西大别和东大别高压/超高压变质地体无法对应。所以,我们认为,桐柏地区超高压岩石的缺乏可能缘于华南陆块的俯冲深度向西变浅,大陆深俯冲作用只局限于大别-苏鲁地区。实际上,桐柏杂岩的出露说明桐柏地区的抬升幅度至少不低于西大别地区。

详细的构造解析和年代学资料表明桐柏-大别-苏鲁造山带的最终构造框架定格于早白垩世的构造伸展事件(Eide et al., 1994; Hacker et al., 1995, 2000; Webb et al., 1999, 2001; Faure et al., 1999, 2003; Ratschbacher et al., 2000; 索书田等, 2001a, b; Wang E et al., 2003; Lin et al., 2009; Li SZ et al., 2009, 2010b),这一事件导致了桐柏-大别高压/超高压变质地体向东的侧向挤出,并使岩石向上抬升约10~15km(Hacker et al., 1995; Webb et al., 1999, 2001; Ratschbacher et al., 2000)。桐柏杂岩中沿造山带方向大规模的花岗岩侵位指示构造伸展始于约140Ma,大量变质岩包体(约占桐柏杂岩的20%)的存在可能暗示他们在花岗岩侵入过程中处于近原地状态。伸展变形的时代已被确定为135~119Ma(Webb et al., 1999, 2001; 崔建军等, 2009b; Cui et al., 未发表数据),也代表桐柏-大别地体向东的挤出时间。从地表形态上看,桐柏-大别地体呈楔形,东宽西窄,这可能与桐柏地区遭受到更强烈的侧向挤压有关。

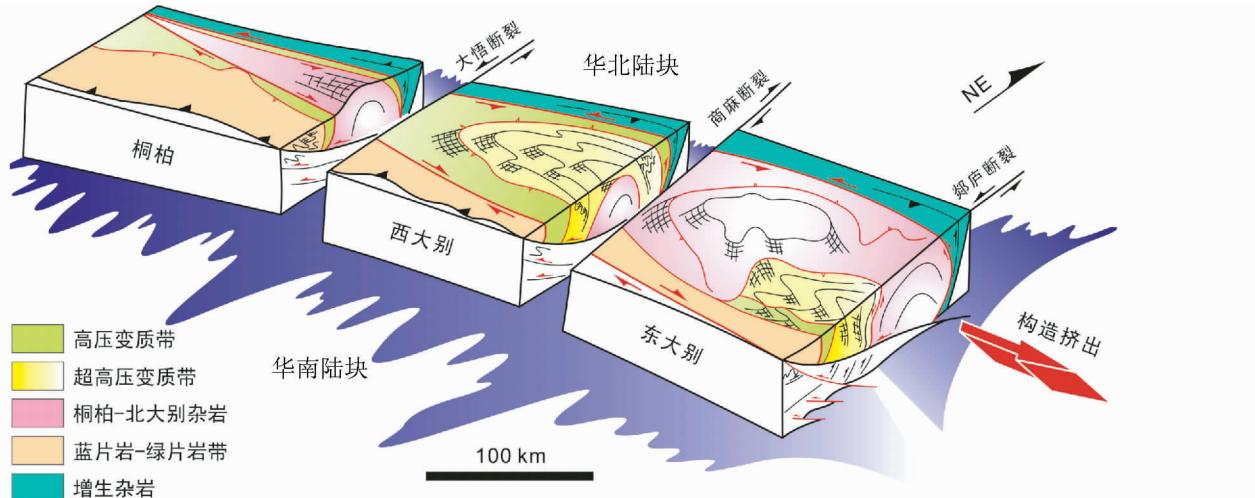


图3 桐柏、西大别和东大别高压/超高压变质地体的地壳结构模型(据 Liu et al. , 2010)

Fig. 3 Interpretative model for the crustal structure of the Tongbai, western Dabie and eastern Dabie HP/UHP metamorphic terranes (after Liu et al. , 2010)

表3 桐柏、西大别、东大别和苏鲁地区(地体)高压/超高压岩片划分及俯冲/折返时代差异

Table 3 Division and age differences of HP/UHP slices in the Tongbai, western Dabie, eastern Dabie and Sulu regions (terrane)

地区(地体)	岩片划分	俯冲(高压/超高压变质)年龄(Ma)	折返(退变质/冷却)年龄(Ma)	数据来源
桐柏	高压岩片 I	255	238	Liu XC et al. (2008, 2010)
	高压岩片 II	232 ~ 220	208 (?)	
西大别	构造混杂岩片		241	Liu XC et al. (2004a, b); Wu et al. (2008); Cheng et al. (2010b)
	低温高压岩片		231	
	高压岩片	243 ~ 238	225 ~ 210	
	超高压岩片	227	213	
东大别	黄镇超高压岩片	242 ~ 236	222	Liu YC et al. (2007a)
	南大别超高压岩片	238 ~ 230	218 ~ 209	
	北大别超高压岩片	219	191	
苏鲁	高压岩片 I	245	231	Liu FL et al. (2009)
	超高压岩片 II	228	215	
	超高压岩片 III	218	202	

## 7 大陆俯冲的起始时间与高压/超高压岩片的穿时俯冲与折返

对西大别地区熊店、胡家湾和苏家河等地的洋壳榴辉岩及其围岩的锆石 U-Pb 定年表明, 华北与华南陆块碰撞之前的大洋岩石圈俯冲发生在约 315Ma (简平等, 2000; Li SG et al. , 2001; Sun et al. , 2002; Cheng et al. , 2009; Wu et al. , 2009; 刘小驰等, 2009; 杨赛红等, 2009), 锆石的微量元素特征以及其中的矿物包裹体类型为这一结论提供了具有说服力的证据。在桐柏地区高压岩片 I 陆壳榴辉岩中获得的两组锆石 U-Pb 年龄、两组全岩-石榴石 Lu-Hf 等时线年龄和一组全岩-石榴石-绿辉石-角闪石-多硅白云母 Rb-Sr 等时线年龄均集中在约 255Ma, 说明华北与华南陆块的碰撞及伴随的陆壳俯冲起始于早二叠世。这一结论得到如下二个证据的

支持。其一, 在熊店及其附近榴辉岩中获得的全岩-石榴石 Lu-Hf 和 Sm-Nd 等时线年龄或石榴石-绿辉石结线年龄集中在 271 ~ 252Ma (Jahn et al. , 2005; Cheng et al. , 2009, 2010a), 明显晚于锆石 U-Pb 年龄, 说明这些洋壳榴辉岩又遭受到陆壳俯冲作用的影响。其二, 在东大别佛子岭群中获得的白云母<sup>40</sup>Ar/<sup>39</sup>Ar 年龄也集中在 271 ~ 261Ma (牛宝贵等, 1994; Faure et al. , 2003; Ratschbacher et al. , 2006), 一般认为佛子岭群代表俯冲带下盘的增生杂岩。所以, 我们认为洋壳的连续俯冲可能跨越了至少约 60Ma, 即从约 315Ma 一直持续到约 255Ma, 而陆壳深俯冲的起始时间将不早于约 255Ma。

另一方面, 桐柏地区高压岩片 I 的白云母<sup>40</sup>Ar/<sup>39</sup>Ar 定年反映高压变质岩石冷却至约 350°C 的时代约为 238Ma, 这一年龄甚至老于桐柏杂岩中变质岩包体(高压岩片 II)的变质时代(232 ~ 220Ma), 说明当高压岩片 II 被俯冲到地壳深处

并经受高压变质时,其上覆的高压岩片 I 已经折返到中上地壳的水平。高压/超高压岩石在俯冲和折返时代上的差别在西大别、东大别和苏鲁地区均已被识别出来(表 3)。在西大别地区,Liu XC *et al.* (2004a, b)针对变质带的空间分布和年代学差别提出了一个高压/超高压岩片穿时俯冲/折返的构造模型,即华南大陆板片在俯冲时末端不断发生破裂,先期俯冲的岩片在后续板片俯冲时将作为上盘连续向浅部折返,不同岩片的折返时间可相差约 30Ma,而超高压岩片的最终折返则归因于华南大陆下地壳的俯冲。在东大别地区,仅超高压变质杂岩就已区分出了 3 个穿时的(或称解耦的)岩片,即黄镇、南大别和北大别超高压岩片(Liu YC *et al.*, 2007a)。在苏鲁地区,形成时代不同的一个高压岩片和两个超高压岩片也已确定(Liu FL *et al.*, 2009),并且提出了大陆板片多重性差异俯冲与折返的动力学模式(许志琴等, 2005; Liu FL *et al.*, 2009)。由此可见,穿时(或差异)俯冲和折返可能是桐柏-大别-苏鲁造山带中高压/超高压岩石得以回返并保存的主要机制。桐柏地区高压变质杂岩的年代学结果为这种构造模型提供了进一步的支持证据,而该区的高压岩片 I 可能代表造山带中最早俯冲和折返的陆壳岩片。

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